

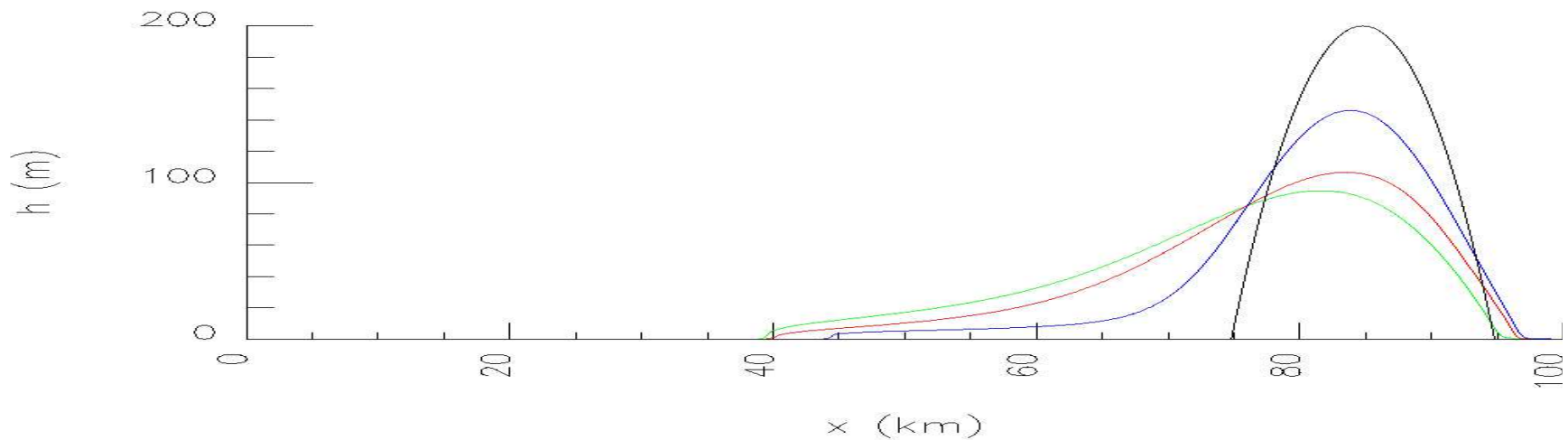
Estimation of Friction Parameters and Laws in 2D Shallow-Water Gravity Currents on the f-Plane, by Data Assimilation



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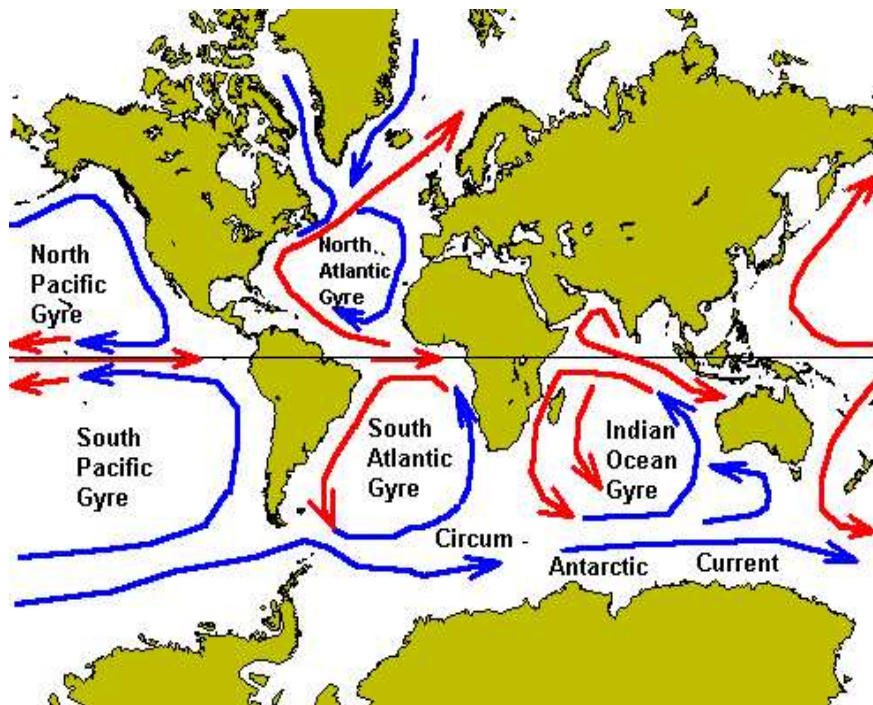
Etudes et paramétrisation
des flux turbulents pour les
courants gravitaires océaniques
par assimilation de données



Ocean Circulation

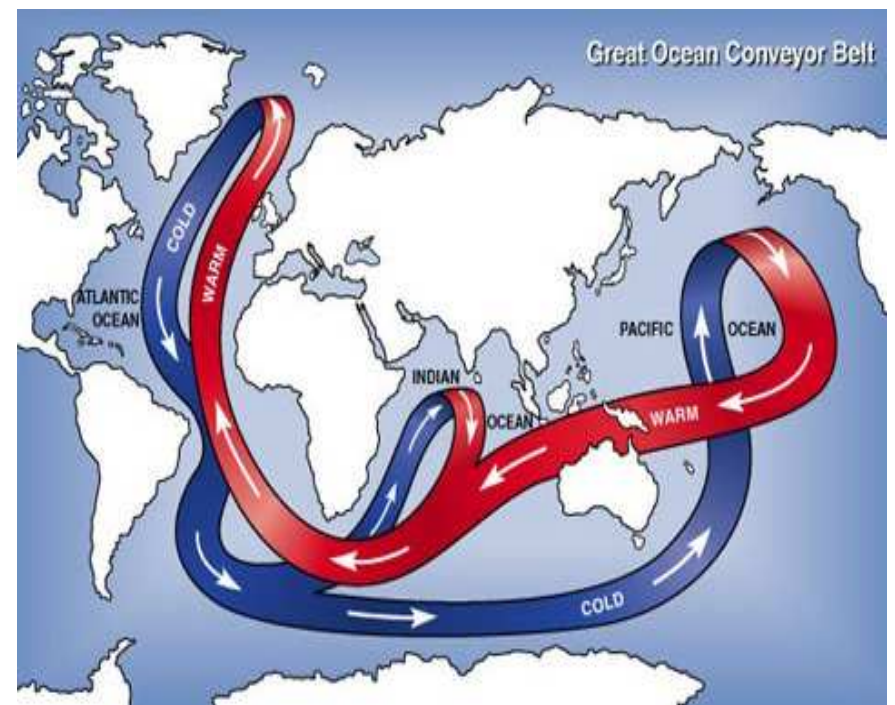
Gyre

“weather”

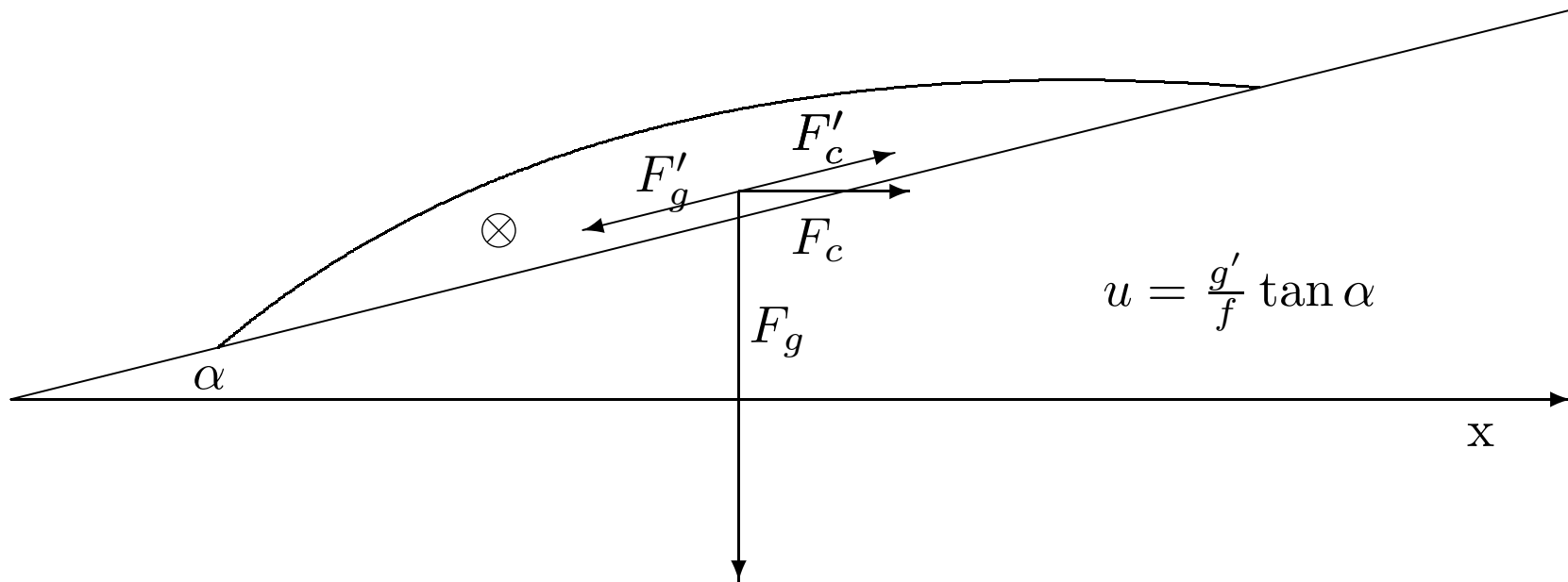
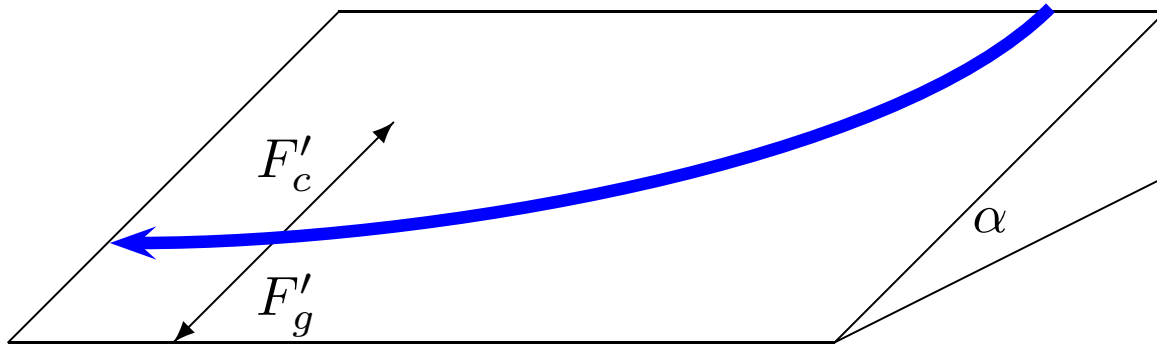


Overturning

“climate”



Gravity Currents in a Rotating Frame

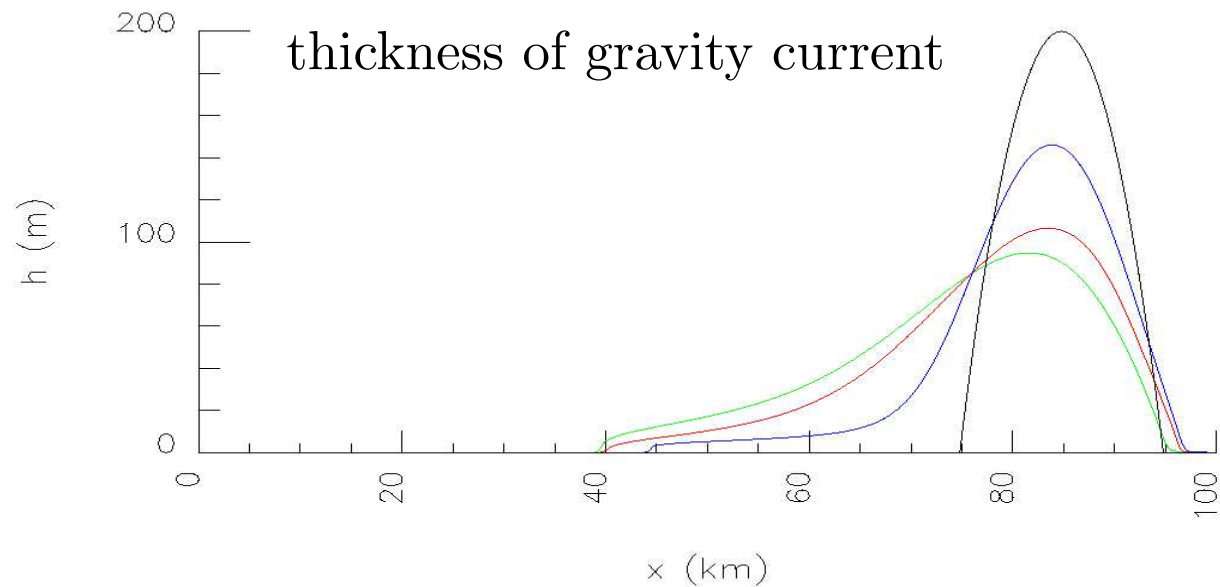


What Makes Gravity Currents Flow Down-Slope ?

- To first order gravity currents change neither depth nor speed (geostrophic equilibrium).
- Dissipative turbulent fluxes of momentum and density perturb the geostrophic equilibrium (friction, mixing, entrainment, detrainment).
- These turbulent fluxes can NOT be explicitly resolved in today's (and tomorrows) models of the global ocean circulation.
- We have to find efficient **parametrisations** and the correct **parameter values**

Time Evolution of a Cross Section

$$\begin{aligned}
 \partial_t u + u \partial_x u - f v + g' (\partial_x h + \tan \alpha) &= -D u + \nu \partial_x^2 u, \\
 \partial_t v + u \partial_x v + f u &= -D v + \nu \partial_x^2 v, \\
 \partial_t h + u \partial_x h + h \partial_x u &= \nu \partial_x^2 h.
 \end{aligned}$$

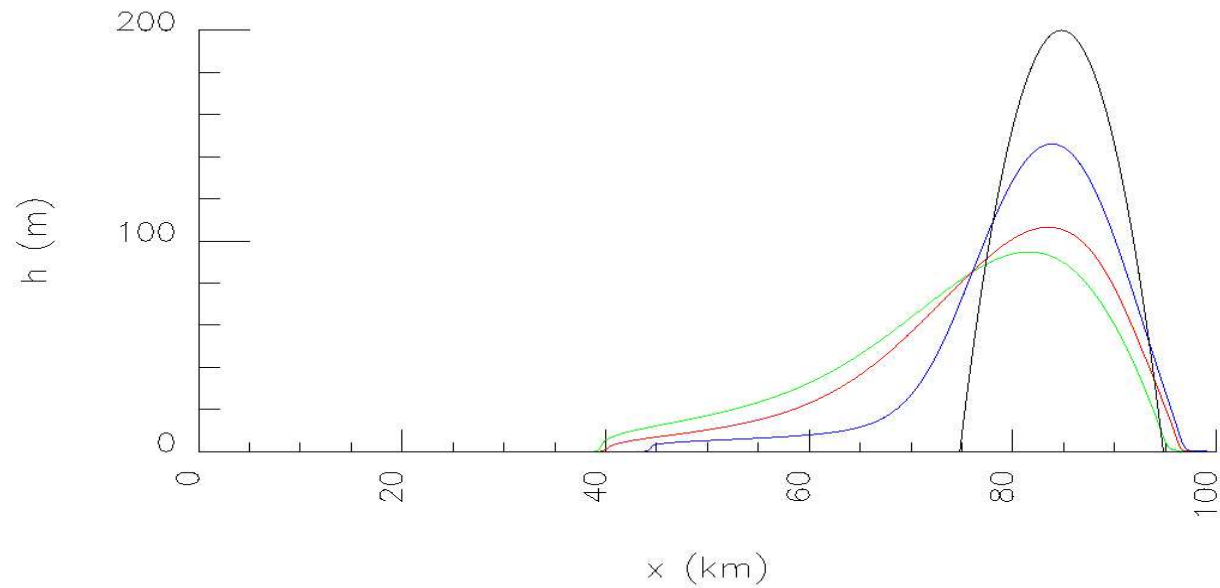


The Friction

$$D = D(x, t) = \frac{1}{h} (\tau + c_D |\mathbf{u}|)$$

Two friction laws:

- Linear Rayleigh friction (τ)
- Quadratic drag law (c_D)



Why use Data Assimilation?

“I am the wisest man alive, for I know one thing, and that is that I know nothing” (Socrates).

I do not know ...

- the roughness of the ocean floor.
- the difference between roughness and obstacles
- the effects of roughness change.
- the roughness type of the ocean floor “k” vrs. “d”.
- the direction of roughness elements.
- the sediment suspension.
- tidal and wave currents (short time-scale)
-

If I knew all this: “The matter is far from being understood” Jiménez, *Ann. Rev. Fluid Mech.* (2004).

The Estimation Procedure

Augmented state vector (containing the parameters):

$$\mathbf{x}(x, t) = (h(x, t), u(x, t), v(x, t), \boldsymbol{\tau}, \mathbf{c}_D) \quad (1)$$

The tool: EnKF (Ensemble Kalman Filter) which updates the ensemble mean and variance by updating each ensemble member:

$$\mathbf{x}_i^a = \mathbf{x}_i^f + \mathbf{K} (h^{obs} + \epsilon_i - \mathbf{H}\mathbf{x}_i^f) \quad (2)$$

$$\mathbf{K} = \mathbf{P}\mathbf{H}^T (\mathbf{H}\mathbf{P}\mathbf{H}^T + \mathbf{W})^{-1} \quad (3)$$

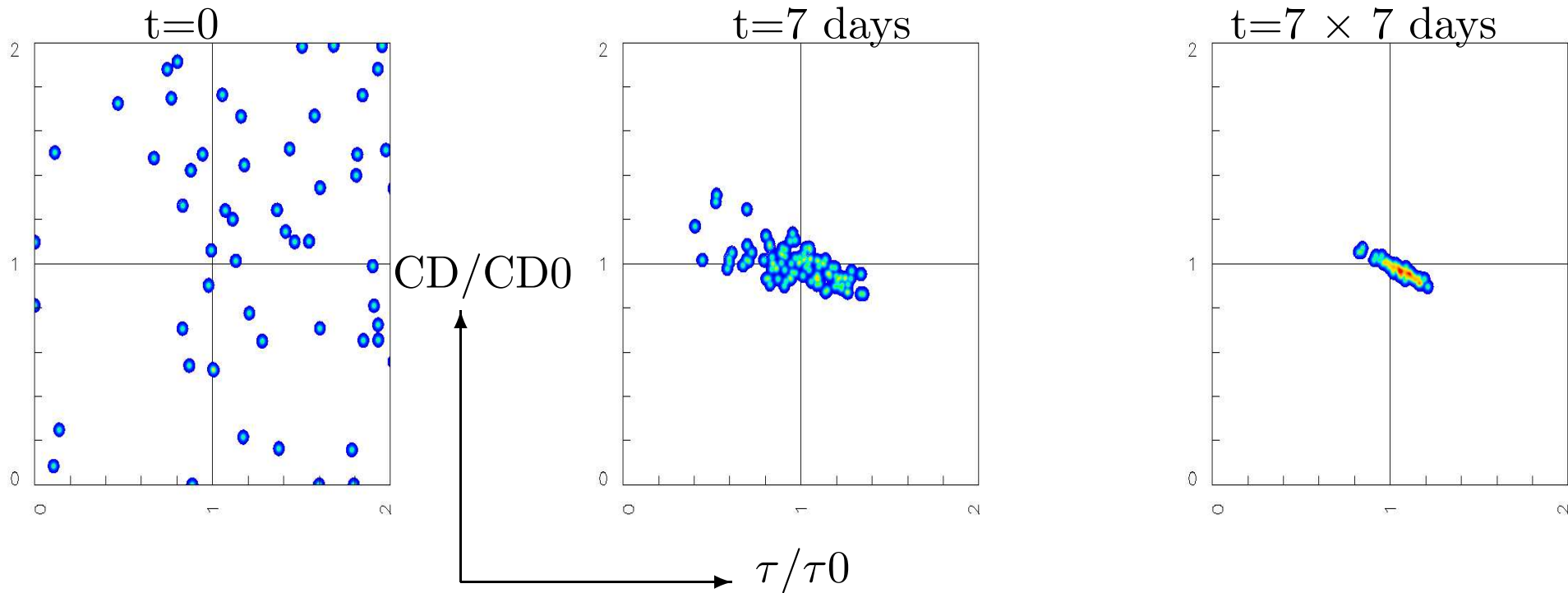
$$\mathbf{P} = \frac{1}{m-1} \sum_{i=1}^m (\mathbf{x}_i^f - \langle \mathbf{x}^f \rangle) (\mathbf{x}_i^f - \langle \mathbf{x}^f \rangle)^T, \quad (4)$$

Choice of initial distribution of parameter values?

Can we obtain r and c_D from ONLY observing $h(x)$?

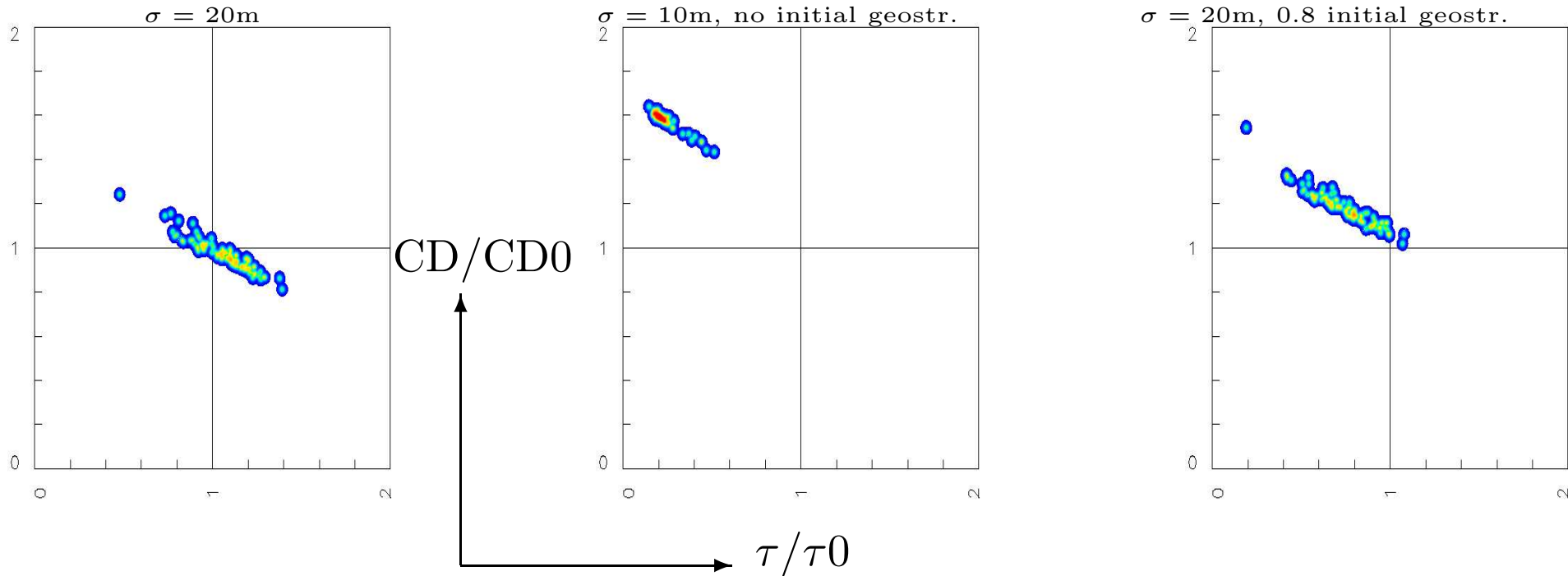
Assimilation every hour, with an observation error of $10m$,

$$\tau_0 = 5 \cdot 10^{-4} \text{ ms}^{-1}, c_{D0} = 5 \cdot 10^{-3}$$



Yes!

Why is the ensemble aligned?



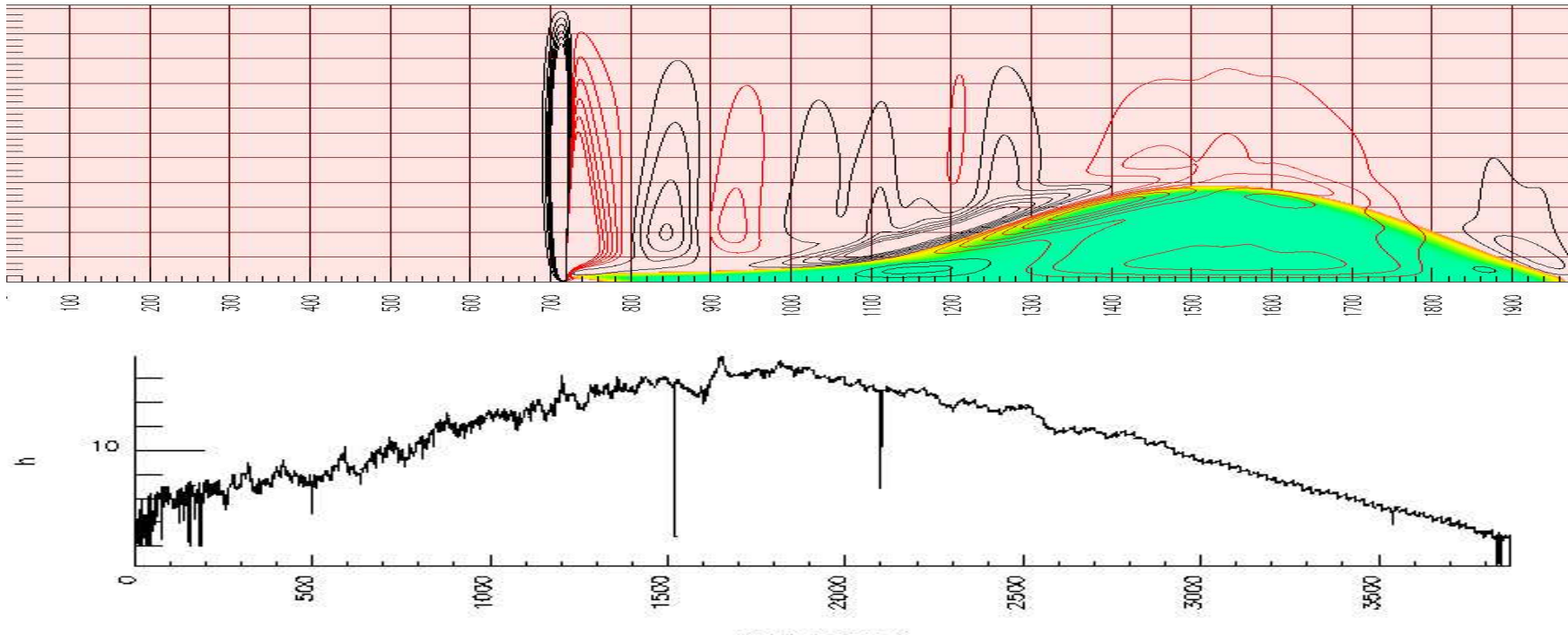
Because it is easy to estimate the average friction $\langle D \rangle_{x,t}$, where:

$$D = (\tau + c_D |\mathbf{u}|)$$

but it is more difficult to separate it into linear Rayleigh friction vs. quadratic drag law. We have a fast projection on the “slow manifold” corresponding to $\langle D \rangle_{x,t} = \text{const.}$ followed by a slow convergence within the “slow manifold.”

Next:

- Move from twin experiments to connecting models in a “lowrarchy”.
- Connecting models to experiments.
- Estimate parameters and determine laws.



Conclusions:

- The EnKF is a robust tool for estimating parameters
- Parameter estimation can be used to determine laws (by estimating parameters).
- The spread of the ensemble tells us something about the physics of the problem.
- The here presented results are likely to be general for parameter estimation.